Wavenumber integration for multilayer ocean environments

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I. SUMMARY

Assume a horizontally stratified environment shown in Fig. 1. The source lies in $(r_s, z_s) = (0, 35)$, the receiver lies in $(r_r, z_r) = (0, 25)$. There are $N_{\text{layer}} = 30$ layers. From Fig. 1, one can see that there are $N_{\text{layer}} - 1 = 29$ interfaces. For the first interface (the layer between one and two), its position is z = 0. Let z_i denote the position (depth) of the bottom of the *i*th layer. In the problem, we set $z_i = (i - 1)d_z$, $i = 1, \dots, N_{\text{layer}} - 1$, where $d_z = 5$. For the top $N_{\text{layer}} - 1$ layers, we set the sound speed $c_i = 1500m/s$, the density $\rho_i = 1g/cm^3$, $i = 1, \dots, N_{\text{layer}} - 1$. For the last layer, we set $z_{N_{\text{layer}}} = N_{\text{layer}}d_z$. And $c_{N_{\text{layer}}} = 1800m/s$, $\rho_{N_{\text{layer}}} = 1.8g/cm^3$. The source strength and the pressure are

$$S_{\omega} = -\frac{4\pi}{\rho\omega^2},\tag{1}$$

$$p(r) = \rho \omega^2 \psi(\omega, r), \qquad (2)$$

where $\psi(\omega, r)$ is the displacement potential.



Fig. 1: Horizontally stratified environment

We define $\mathbf{v}_m(k_r, z_m)$

$$\mathbf{v}_m(k_r, z_m) = \begin{bmatrix} w(k_r, z) \\ \sigma_{zz}(k_r, z) \end{bmatrix}, \quad \text{the bottom of the m th layer or the top of the m + 1 th layer}$$
(3)

which contains the displacements and stresses at interface m. Similarly, define the degree-of-freedom vector $\mathbf{a}_m(k_r)$ for layer number m as composed of the complex amplitudes of the downgoing and upgoing solutions

$$\mathbf{a}_m(k_r) = \begin{bmatrix} A_m^-(k_r) \\ A_m^+(k_r) \end{bmatrix},\tag{4}$$

define the local coefficient matrix $\mathbf{c}_m(k_r,z)$

$$\mathbf{c}_m(k_r, z) = \begin{bmatrix} -\mathrm{i}k_z \mathrm{e}^{-\mathrm{i}k_z z} & \mathrm{i}k_z \mathrm{e}^{\mathrm{i}k_z z} \\ -\rho_m \omega^2 \mathrm{e}^{-\mathrm{i}k_z z} & -\rho_m \omega^2 \mathrm{e}^{\mathrm{i}k_z z} \end{bmatrix}.$$
(5)

We know that

$$\mathbf{v}_m(k_r, z_m) = \mathbf{c}_m(k_r, z_m) \mathbf{a}_m(k_r), \quad \mathbf{v}_m(k_r, z_{m-1}) = \mathbf{c}_m(k_r, z_{m-1}) \mathbf{a}_m(k_r).$$
(6)

The above relationship reveals that

$$\mathbf{v}_m(k_r, z_{m-1}) = \mathbf{P}_m(k_r) \mathbf{v}_m(k_r, z_m) = \mathbf{c}_m(k_r, z_{m-1}) [\mathbf{c}_m(k_r, z_m)]^{-1} \mathbf{v}_m(k_r, z_m).$$
(7)

Using the continuity of the field parameters at the interfaces, we can establish a matrix relation between the field parameters at some interface m and the parameters at a lower interface n (n < m)

$$\mathbf{v}_{n}(k_{r}, z_{n}) = \mathbf{P}_{n+1}(k_{r})\mathbf{v}_{n+1}(k_{r}, z_{n+1}) = \prod_{i=n+1}^{m} \mathbf{P}_{i}(k_{r})\mathbf{v}_{m}(k_{r}, z_{m}).$$
(8)

Since

$$\sigma_{zz}(k_r, z_{N-1}) = -\frac{\rho_N \omega^2}{ik_{z,N}} w(k_r, z_{N-1}),$$
(9)

we rewrite $\mathbf{v}_{N-1}(k_r, z_{N-1})$ as

$$\mathbf{v}_{N-1}(k_r, z_{N-1}) = \begin{bmatrix} w(k_r, z_{N-1}) \\ \sigma_{zz}(k_r, z_{N-1}) \end{bmatrix} = \begin{bmatrix} 1 \\ -\frac{\rho_N \omega^2}{\mathbf{i} k_{z,N}} \end{bmatrix} w(k_r, z_{N-1}),$$
(10)

DRAFT

Set $w(k_r, z_{N-1}) = 1$ (we will solve $w(k_r, z_{N-1})$ later), and calculate $\mathbf{v}_r(k_r, z_r)$ and $\mathbf{v}_s(k_r, z_s)$ at the receiver and source interface.

Calculate the uppermost surface $\mathbf{v}_1(k_r,z_1)$ as

where

$$\hat{\mathbf{v}}(k_r, z_s) = \begin{bmatrix} -\frac{S_\omega}{2\pi} \\ 0 \end{bmatrix}.$$
(12)

We obtain

$$w(k_r, z_{N-1}) = -\frac{\hat{\sigma}_{zz}(k_r, z_1)}{\sigma_{zz}(k_r, z_1)}$$
(13)

via (11). Knowing $\mathbf{v}(k_r, z)$, using the Hankel transform for each component of $\mathbf{v}(k_r, z)$

$$f(r,z) = \int_0^\infty f(k_r, z) J_0(k_r r) k_r \mathrm{d}k_r, \tag{14}$$

we obtain $\mathbf{v}(r, z)$.

REFERENCES

[1] "Computational Ocean Accoustics,"